

# A 3D baroclinic model of the Burdekin River Plume, Australia

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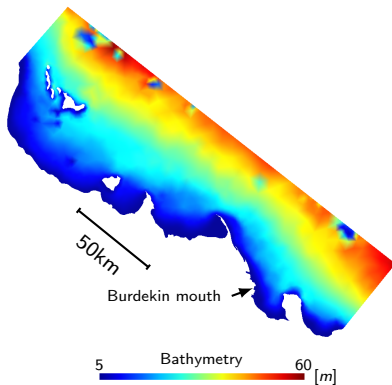
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## The problem

*Goal* : Understanding the key processes controlling the fate of sediment exported by the Burdekin River to the Great Barrier Reef [Lewis et al., 2014]



⇒ 3D modelling

# SLIM 3D : a baroclinic dg-finite element model

## Second-generation Louvain-la-Neuve Ice-ocean Model<sup>1</sup>

- Spatial features
  - $P_1$  dg-finite element discretisation
  - prismatic elements
  - ALE formulation on a moving mesh
- Time discretisation
  - split-explicit approach
  - implicit vertical diffusion on every column

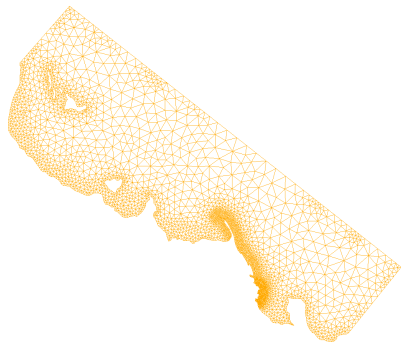
[Kärnä et al., 2013]

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<sup>1</sup>[www.climate.be/slim](http://www.climate.be/slim)

# Applying SLIM 3D to the Burdekin River Plume dynamics

- Until now:
  - Square boxes geometry
  - flat or linear bathymetry
- The Challenge:
  - Complex geometry
  - Complex bathymetry
  - Actual forcings

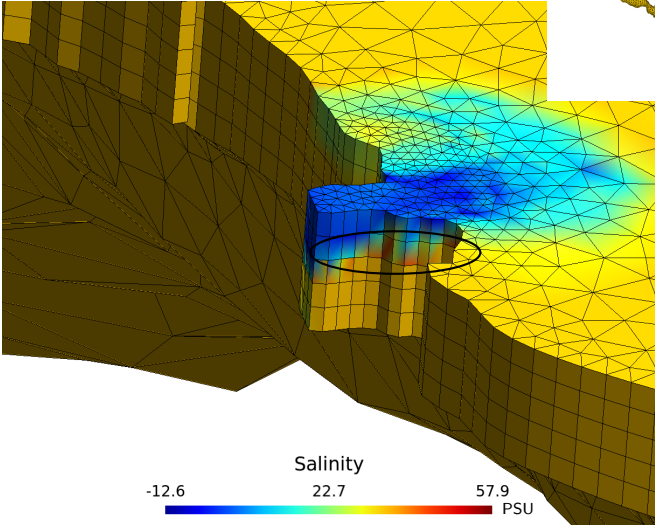


Mesh generated with Gmsh Software<sup>2</sup>

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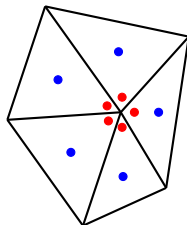
<sup>2</sup>[www.geuz.org/gmsh](http://www.geuz.org/gmsh)

# Overshoot problems



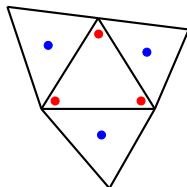
## Limiters are necessary !

- Kuzmin [2010], Aizinger [2011]



- dg nodes
- P<sub>0</sub> values

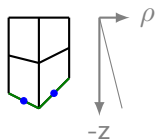
- Cockburn and Shu [1998]



## Limiter issues

No more overshoots, but it has a cost:

- Invalid lake at rest for a stratified water column.  
⇒ Choose boundary condition for limiter



- Constant fields on an element where strong gradients appear  
⇒ Loss of precision. ⇒ Adaptive mesh
- This can lead to non-physical behaviour for  $\sigma$ -layers.  
⇒ Constant depth for upper layers

## Moving mesh difficulties

- ALE formulation:

$$\frac{\partial T}{\partial t} + T \frac{\partial w_m}{\partial z} + \nabla_h \cdot (\mathbf{u}T) + \frac{\partial((w - w_m)T)}{\partial z} = \mathbf{D}$$

- Moving mesh algorithm:

$$z = -h + (h + \eta) \frac{z_0 + h}{h}, \quad z_0 \in [-h, 0], z \in [-h, \eta]$$

$$\Rightarrow \begin{cases} w_m &= w_{m,\text{surf}} \frac{z_0 + h}{h} \\ \frac{\partial w_m}{\partial z} &= \frac{w_{m,\text{surf}}}{h + \eta} \end{cases}$$

- $w_m$  and  $\partial w_m / \partial z$  must match perfectly !



## Moving mesh difficulties (2)

$$z = -h + (h + \eta) \frac{z_0 + h}{h}$$

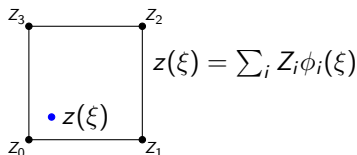
- $\sigma$ -layers mesh:

$$z_0 = -\alpha h, \quad \alpha \in [0, 1]$$
$$\Rightarrow \begin{cases} w_m &= w_{m,\text{surf}} \frac{z_0 + h}{h} = w_{m,\text{surf}}(1 - \alpha) \\ \frac{\partial w_m}{\partial z} &= \frac{w_{m,\text{surf}}}{h + \eta} \end{cases}$$

- Constant depth for upper layers:

$$\Rightarrow \begin{cases} w_m &= w_{m,\text{surf}} \frac{z_0 + h}{h} = w_{m,\text{surf}} \frac{z + h}{h + \eta} \\ \frac{\partial w_m}{\partial z} &= \frac{w_{m,\text{surf}}}{h + \eta} \end{cases}$$

$$\frac{\partial}{\partial z} (w_m) = w_{m,\text{surf}} \frac{1}{h} \frac{\partial z_0}{\partial z} \not\approx \frac{w_{m,\text{surf}}}{h + \eta}$$



# Upper layers at constant depth

## Salinity Profile

- 5  $\sigma$ -layers mesh

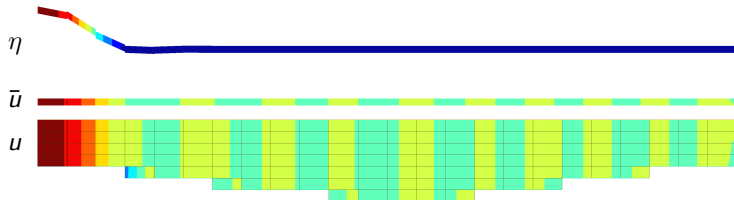


- 7 layers mesh: first two at constant depth, next 5  $\sigma$ -layers

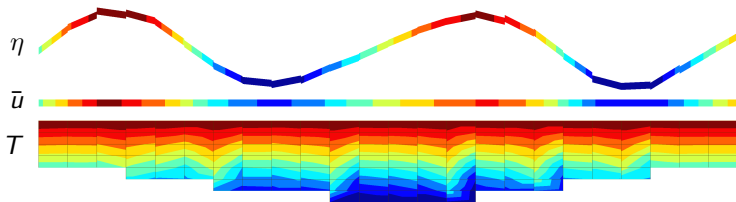


## When z layers are needed..

- Shift on the entire column

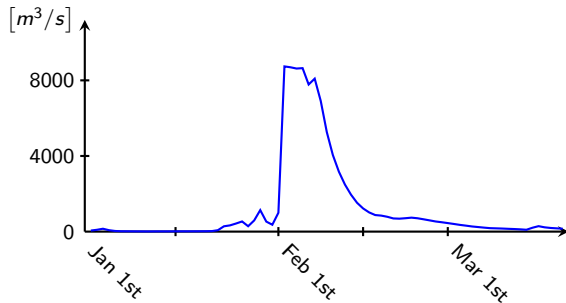


- No shift where vertical bottom boundary



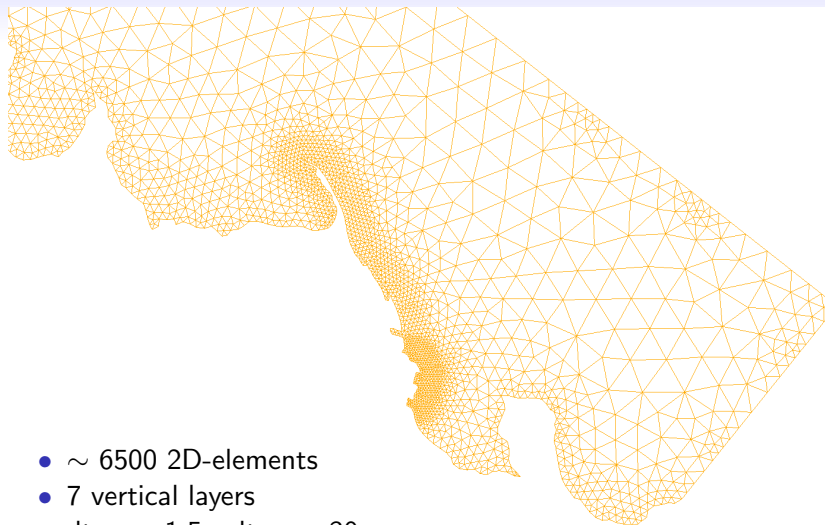
# Burdekin River Plume dynamics

- Burdekin river discharge for 2007 flood season



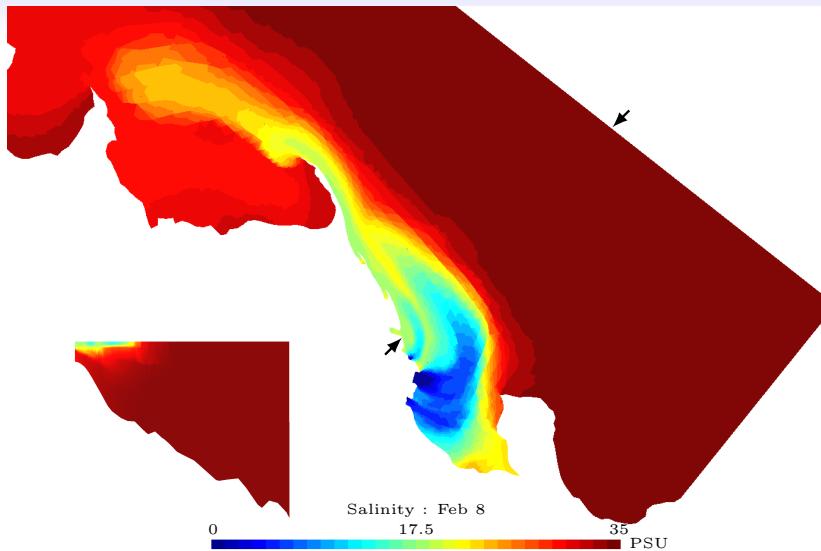
- Wind forcing
- Tidal forcing
- Varying sediment concentration discharge

## Salinity dynamics



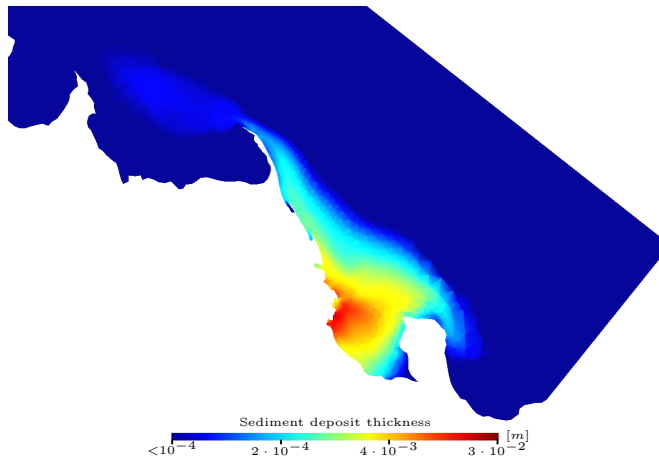
- $\sim 6500$  2D-elements
- 7 vertical layers
- $dt_{2D} = 1.5s$ ,  $dt_{3D} = 30s$
- 20 times faster than physical time on 12 CPUs

## Salinity dynamics



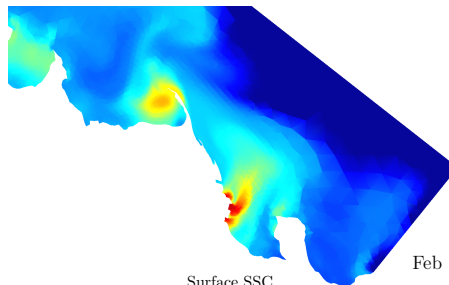
## Results

- Only with settling effect :  
>50km region : deposit thickness < 0.1mm  
⇒ accordance with Lewis et al. [2014]



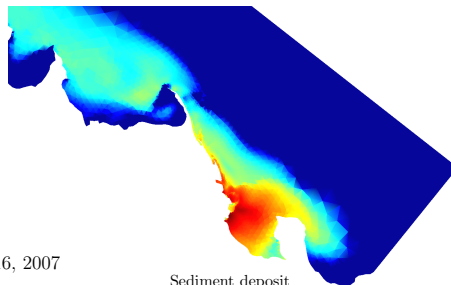
## Results (2)

- Full sediment model :  
>50km region : deposit thickness < 0.1mm  
⇒ accordance with Lewis et al. [2014]



Surface SSC  
 $\text{kg/m}^3$   
<  $10^{-4}$     $7 \cdot 10^{-4}$     $5 \cdot 10^{-3}$     $4 \cdot 10^{-2}$     $3 \cdot 10^{-1}$

Feb 16, 2007



Sediment deposit  
m  
<  $10^{-5}$     $7 \cdot 10^{-5}$     $5 \cdot 10^{-4}$     $4 \cdot 10^{-3}$     $3 \cdot 10^{-2}$



## Conclusions and Perspective

- Burdekin River model
  - During a flood event, the sediments do stay in the region defined by Lewis et al. [2014].
  - Calibration of sediment model needs to be done.
- SLIM 3D model
  - SLIM 3D is ready to model actual coastal applications for a small ratio  $\max(\text{bath})/\min(\text{bath})$ .
  - More flexibility for vertical layers should be developed.

Thank you for your attention !

## References

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